

1  
2 Carbon consequences of forest disturbance and recovery across the conterminous  
3 United States

4  
5 **Authors:** Christopher A. Williams<sup>1,\*</sup>, G. James Collatz<sup>2</sup>, Jeffrey Masek<sup>2</sup>, Samuel N. Goward<sup>3</sup>

6 <sup>1</sup>Graduate School of Geography, Clark University, Worcester, MA, USA

7 <sup>2</sup>NASA Goddard Space Flight Center - Code 618, Biospheric Sciences Laboratory, Greenbelt,  
8 MD, USA

9 <sup>3</sup>Department of Geography, University of Maryland, College Park, MD, USA

10 \*Corresponding Author: [cwilliams@clarku.edu](mailto:cwilliams@clarku.edu), P: 508-793-7323; F: 508-793-8881

11  
12 **Second Revision for:** Global Biogeochemical Cycles  
13  
14 **Running title:** Forest disturbance and carbon dynamics  
15  
16 **Keywords:** net ecosystem productivity; carbon sequestration; forest inventory and analysis;  
17 carbon cycle modeling; forest carbon sink attribution  
18  
19  
20

21   **Abstract:**

22   Forests of North America are thought to constitute a significant long term sink for atmospheric  
23   carbon. The United States Forest Service Forest Inventory and Analysis (FIA) program has  
24   developed a large data base of stock changes derived from consecutive estimates of growing stock  
25   volume in the US. These data reveal a large and relatively stable increase in forest carbon stocks  
26   over the last two decades or more. The mechanisms underlying this national increase in forest  
27   stocks may include recovery of forests from past disturbances, net increases in forest area, and  
28   growth enhancement driven by climate or fertilization by CO<sub>2</sub> and Nitrogen. Here we estimate  
29   the forest recovery component of the observed stock changes using FIA data on the age structure  
30   of US forests and carbon stocks as a function of age. The latter are used to parameterize forest  
31   disturbance and recovery processes in a carbon cycle model. We then apply resulting  
32   disturbance/recovery dynamics to landscapes and regions based on the forest age distributions.  
33   The analysis centers on 28 representative climate settings spread about forested regions of the  
34   conterminous US. We estimate carbon fluxes for each region and propagate uncertainties in  
35   calibration data through to the predicted fluxes. The largest recovery-driven carbon sinks are  
36   found in the Southcentral, Pacific Northwest, and Pacific Southwest regions, with spatially  
37   averaged net ecosystem productivity (*NEP*) of about 100 g C m<sup>-2</sup> a<sup>-1</sup> driven by forest age  
38   structure. Carbon sinks from recovery in the Northeast and Northern Lake States remain  
39   moderate to large owing to the legacy of historical clearing and relatively low modern disturbance  
40   rates from harvest and fire. At the continental scale, we find a conterminous U.S. forest *NEP* of  
41   only 0.16 Pg C a<sup>-1</sup> from age structure in 2005, or only 0.047 Pg C a<sup>-1</sup> of forest stock change after  
42   accounting for fire emissions and harvest transfers. Recent estimates of *NEP* derived from  
43   inventory stock change, harvest, and fire data show twice the *NEP* sink we derive from forest age

44 distributions. We discuss possible reasons for the discrepancies including modeling errors and  
45 the possibility of climate and/or fertilization ( $\text{CO}_2$  or N) growth enhancements.

46

47    **1. Introduction**

48    The global imbalance among ocean, industrial, and land use sources/sinks of CO<sub>2</sub> and the  
49    amount accumulating in the atmosphere implies significant net CO<sub>2</sub> uptake by the terrestrial  
50    biosphere [e.g. *Schimel et al.*, 2001; *Tans et al.*, 1990]. Despite large uncertainty about  
51    magnitude and process, analyses tend to point to northern temperate and boreal lands as dominant  
52    terrestrial sinks of CO<sub>2</sub> but with considerable controversy regarding attribution to specific regions  
53    or continents [e.g. *Bousquet et al.*, 2000; *Fan et al.*, 1998; *Gurney et al.*, 2002; *Kaminski et al.*,  
54    2001; *Myneni et al.*, 2001; *Tans et al.*, 1990]. However, some recent work suggests far smaller  
55    sinks in northern temperate and boreal lands [*Ito et al.*, 2008; *Stephens et al.*, 2007; *Yang et al.*,  
56    2007].

57    Estimates of the conterminous U. S. forest net carbon uptake from the atmosphere range from  
58    only 10 to over 200 Tg C a<sup>-1</sup> [*EPA*, 2010; *Houghton et al.*, 1999; *King et al.*, 2007; *Pacala et al.*,  
59    2001; *Turner et al.*, 1995] in the last 2 decades. Note that here we consider the forest stock  
60    change alone rather than the forest sector stock change, where the latter also includes carbon  
61    accumulated in wood products (see State of the Carbon Cycle Report [*King et al.*, 2007]).

62    Techniques for estimating forest carbon fluxes at regional to national scales include three  
63    approaches. The stock change method is exemplified in the US report to the United Nations  
64    Framework Convention for Climate Change [e.g. *EPA*, 2008] which uses US Forest Service  
65    Forest Inventory and Analysis (FIA) data on sequential measurement of tree diameters and/or  
66    wood volumes for about 100,000 forest plots at 5-20 year intervals. Allometric and biomass  
67    expansion factors are used to convert volume into forest carbon stocks. The rate of carbon uptake

68 is then estimated as the difference between sequential measurements divided by the number of  
69 years in the interval.

70 Another technique for estimating forest carbon sinks combines estimates of the stand age  
71 structure of forests with age-specific carbon accumulation rates, termed the “age-accumulation”  
72 approach in this work. These carbon accumulation rates are inferred from carbon stocks as a  
73 function of age [e.g. *Houghton*, 1999], known as yield tables in forestry literature, and may be  
74 derived empirically from inventory estimates of stand volume and age or from a process oriented  
75 dynamic growth model. Finally, forest carbon sinks have been estimated from process models  
76 that account for the effects of climate variability and CO<sub>2</sub> and nitrogen fertilization but not  
77 necessarily for land use and disturbance processes [e.g. *Schimel et al.*, 2000]. These effects are  
78 fully contained in the stock change method because it relies on contemporary changes in stocks,  
79 but the age-accumulation approach relies on a historical characterization of carbon stock  
80 accumulation and thus misses some of the contemporary influences (see Auxiliary 4).

81 Forest stock changes result from the sum of net ecosystem productivity (*NEP*), fire losses, and  
82 harvest (see Figure 1). Significant decreases in harvest and fire have not been observed over the  
83 past few decades so speculation as to the mechanisms underlying the stock increases have focused  
84 more on growth enhancement from either climate change or fertilization with elevated carbon  
85 dioxide or nitrogen [*Houghton*, 1999; *McGuire et al.*, 2001; *Nemani et al.*, 2002; *Pan et al.*, 2009;  
86 *Schimel et al.*, 2000; *Zhou et al.*, 2003] and on forest growth from post-disturbance recovery or  
87 fire suppression [*Caspersen et al.*, 2000; *Hurtt et al.*, 2002; *Pacala et al.*, 2001]. Though the  
88 growth enhancement hypothesis has been challenged by *Casperon et al.* [2000] using forest  
89 inventory data, others have argued that plausible rates of growth enhancement cannot be detected  
90 using existing inventories [*Joos et al.*, 2002] and recent work presents observational evidence

91 supporting a large climate change or fertilization induced sink [Cole *et al.*, 2010; McMahon *et al.*,  
92 2010; Thomas *et al.*, 2009].

93 Disturbed forests, if not converted to another land cover type, have the potential to regrow,  
94 recover, or even surpass pre-disturbance carbon stocks over decades to several hundred years.

95 The long-standing dogma of the carbon source/sink dynamics for stand-replacing disturbance  
96 involves a rapid pulse emission followed by sizeable net uptake that gradually declines [Koerner,  
97 2003; Odum, 1969]. This pattern is broadly supported by chronosequence observations of carbon  
98 stocks [Bond-Lamberty *et al.*, 2004; Gough *et al.*, 2007; Pregitzer and Euskirchen, 2004; Richter  
99 *et al.*, 1999; Thornton *et al.*, 2002] and forest-atmosphere net CO<sub>2</sub> exchange [Amiro *et al.*, 2011;  
100 Barford *et al.*, 2001; Goulden *et al.*, 2011; Law *et al.*, 2003; Schwalm *et al.*, 2007], but the precise  
101 post-disturbance carbon dynamics vary by forest type and climate and this detail remains poorly  
102 characterized.

103 The analysis reported here attempts comprehensive assessment of the carbon consequences of  
104 past and present forest disturbance and recovery across the conterminous United States. We ask  
105 if the forest age structure of the conterminous US forests accounts for the stock changes reported  
106 by the FIA. Our approach utilizes the national forest inventory data (and uncertainties) to  
107 constrain the forest disturbance and recovery processes represented in an ecosystem carbon cycle  
108 model to obtain regional and national estimates of carbon consequences. The basic method can  
109 be described as having two main steps. First, we derive forest type and climate specific post-  
110 disturbance NEP trajectories by fitting a first-order terrestrial carbon cycle model (CASA, [Potter  
111 *et al.*, 1993; Randerson *et al.*, 1996]) to grow wood stocks consistent with FIA data. Second,  
112 these characteristic trajectories are applied to landscapes with forest age maps obtained from FIA  
113 age distributions to derive maps of NEP and biomass. As such, our approach corresponds to the

114 age-accumulation method for estimating forest carbon sinks as described above. Results  
115 represent carbon dynamics of forested ecoregions across the conterminous US to provide a  
116 continental-scale view of forest recovery from past disturbances. In addition, we formally  
117 propagate the uncertainty in FIA age-biomass trends using a Monte Carlo approach, as well as  
118 examine to what degree results are sensitive to uncertainty in the model's parameterization of  
119 carbon turnover time, and dependence on light, moisture, and temperature. Discrepancies  
120 between FIA estimates of stock changes and those from our age-accumulation modeling are  
121 assessed in terms of modeling errors and potential growth enhancements above and beyond  
122 recovery, similar to *Houghton* [2003].

123

## 124 **2. Methods**

### 125 **2.1 Overview**

126 The core of our approach is to estimate the frequency ( $F$ ) of land area in a region ( $A_{reg}$ ), as  
127 well as the flux or stock of carbon ( $Q$ ) each within strata of stand age, forest type (e.g. Aspen-  
128 Birch), and site productivity (high or low) (denoted with  $a, f, p$  subscripts). Regions are defined  
129 according to the Resource Planning Act Assessment by the US Forest Service. From this we  
130 calculate the regional mass flux or stock ( $Q_{reg,s}$ ) for a particular climate setting (subscript  $s$ ) within  
131 each region, as well as its uncertainty ( $\delta$ , described further below), according to

132

$$133 Q_{reg,s} = \sum_a \sum_f \sum_p Q_{afp} F_{afp} A_{reg} \quad (1),$$

134

135

136 where  $F$  is the frequency of forest area adjusted to sum to unity over the three strata and obtained  
 137 from the regional FIA samples of the area of forest land as described in Section 2.2,  $A_{reg}$  is the  
 138 total forested area in the region, and subscripts are:  $a$  for stand age,  $f$  for forest type group, and  $p$   
 139 for productivity class. The work reported here is part of a larger project to incorporate stand age  
 140 derived from Landsat time series data. In this parallel effort, specific scenes for Landsat time  
 141 series were obtained from a statistically rigorous sampling procedure of forest type spatially  
 142 dispersed within Eastern and Western regions [Goward *et al.*, 2008]. Here we use the climate  
 143 (temperature, precipitation, incident solar radiation) and phenology for each scene (Figure 2) to  
 144 simulate fluxes and stocks for each forest type and productivity class within the scene. The scenes  
 145 within a region are generally good representations of the region except for the Pacific Southwest  
 146 where coastal forests are not well represented. The scene level fluxes are then aggregated to  
 147 regional forest fluxes and stocks by averaging across the number of climate settings (scenes,  $N_s$ )  
 148 in a region as

149

$$\begin{aligned}
 150 \quad Q_{reg} &= \frac{1}{N_s} \sum_s Q_{reg,s} \\
 &\delta Q_{reg} = \frac{1}{N_s} \sum_s \delta Q_{reg,s}
 \end{aligned} \tag{2},$$

151

152 and conterminous US estimates (subscript  $nat$ ) are obtained from the sum over regions  
 153

$$\begin{aligned}
 154 \quad Q_{nat} &= \sum_{reg} Q_{reg} \\
 &\delta Q_{nat} = \sum_{reg} \delta Q_{reg}
 \end{aligned} \tag{3}.$$

155

156 We note that our estimates do not account for possible changes in forest carbon due to changes in  
157 forest area, though in the Discussion section we explain why this is unlikely to contribute a large  
158 carbon source or sink given the rates of current-day net land conversion.

159 The relationship between fluxes and stocks can be diagrammed as shown in Figure 1. The so-  
160 called forest sector sources/sinks refer to the net flux between the atmosphere and forest stocks  
161 plus wood products stocks. The inventory approach to calculating the net forest-atmosphere flux  
162 involves a measured change in carbon stocks over a specified period. A change in forest carbon  
163 stocks can occur because of changes in the physiological fluxes of photosynthesis and ecosystem  
164 respiration (balanced as *NEP*), as well as changes in disturbance for example by fire or harvest.  
165 *NEP* can then be inferred as the difference between  $\Delta C_{stocks}$  and removals from fire and harvest.  
166 The net forest sector flux to the atmosphere is the sum of  $\Delta C_{stocks}$  and  $\Delta C_{wood\ products}$ . This  
167 approach, used in national reports to United Nations Framework Convention on Climate Change,  
168 derives  $\Delta C_{wood\ products}$  from independent harvest records and empirical decay constants for wood  
169 products and landfills.

170 Our approach is to calibrate our modeled biomass as a function of age using forest inventory  
171 data. We then apply the biomass and associated *NEP* from forest disturbance and recovery to the  
172 landscape based on the forest area reported by the FIA within strata of age, forest types and  
173 productivity classes within each region. In our modeling framework an important driver of  
174  $\Delta C_{stocks}$  is net primary production (*NPP*), and the turnover times of wood and detrital pools. *NPP*  
175 allocated to leaves and fine roots is quickly decomposed and cannot represent a persistent (>  
176 decadal) sink. The turnover rates of wood and its immediate detrital pool, coarse woody debris,  
177 are much slower, on the order of decades, and thus able to account for long-term net carbon fluxes  
178 (on the order of a century). Fluxes from large stocks of slowly overturning soil pools are also

179 slow to respond to disturbance. By the time these large soil pools are affected by disturbance,  
180 recovery may have already occurred. This phenomenon is expressed as a low sensitivity of  $NEP$   
181 to the slow turnover pools in recovering forests (see Auxiliary Material, Auxiliary 1). Of course  
182 the slow soil pools are a significant source or sink in conditions where changes in fluxes into the  
183 slow pools are large and longer term such as in permanent conversion from or to forest. This  
184 approach allows us to map  $NEP$  from recovery, one of the key atmospheric flux components  
185 needed to understand source/sink processes.  $NEP$  is a purely biological flux dependent on  
186 photosynthesis and respiration alone. Fluxes out of the forest arising from harvest or fire combine  
187 with  $NEP$  to produce net biome productivity ( $NBP$ ) which is equivalent to  $\Delta C_{stocks}$ . Note that we  
188 have neglected the generally smaller fluxes that contribute to  $NBP$  such as lateral fluxes of  
189 carbonate and organic matter in liquid form as well as volatile organic carbon emissions (see  
190 *Chapin III et al.* [2006]).

191

## 192 **2.2 Data Sources and Modeling**

193 Flux trajectories are derived by fitting forest growth, mortality and shedding, and allocation  
194 parameters within the Carnegie-Ames-Stanford Approach (CASA) carbon-cycle process model  
195 [*Potter et al.*, 1993; *Randerson et al.*, 1996] to accumulate carbon in aboveground wood biomass  
196 consistent with forest inventory data. Productivity in CASA is represented with a light use  
197 efficiency approach in which  $NPP$  is proportional to the fractional absorption of  
198 photosynthetically active radiation ( $f_{PAR}$ ) times an efficiency term modulated by environmental  
199 conditions.  $NPP$  is allocated to leaves, roots, and wood which have specific turnover rates that  
200 reflect the delivery of carbon to nine detrital pools on the surface and in the soil. These pools  
201 decompose at specific turnover rates that are also modulated by environmental conditions.

202 Disturbance causes *NPP* to initially decrease, and removes or transfers carbon between live and  
203 detrital pools, the atmosphere, and forest harvest. In this implementation, we adjust the default  
204 rate of productivity to match carbon accumulation observed in age-accumulation trajectories from  
205 forest inventory data.

206 Inventory data were obtained from the FIA field plots (FIA Database Version 4), providing  
207 means and sampling errors for two attributes: 1) all live, oven-dry aboveground wood biomass,  
208 and 2) area of forest land. The quotient of these attributes provides biomass per unit area. Each  
209 attribute was sampled within strata of forest type group (28 classes), age (20 year age classes to  
210 200+ years), and lumped into high and low productivity classes, defined as 120 to >225 cubic feet  
211 acre<sup>-1</sup> annum<sup>-1</sup> and 20 to <120 cubic feet acre<sup>-1</sup> annum<sup>-1</sup> respectively. Inventory samples were  
212 drawn for regions defined by the Resource Planning Act Assessment by the US Forest Service  
213 that divides the conterminous U.S. into the Northeast (NE), Southeast (SE), Northern Lakes States  
214 (NLS), South Central (SC), Northern Prairie States (NPS), Rocky Mountain North (RMN), Rocky  
215 Mountain South (RMS), Pacific Southwest (PSW), and Pacific Northwest (PNW) region (Figure  
216 2). FIA data on forest carbon and area that are available via World Wide Web download include  
217 variances for each. However these variances cannot be exactly combined to estimate uncertainty  
218 because of unknown covariance between carbon stock and area [Bechtold and Patterson, 2005].  
219 Statisticians from the FIA (Charles Scott and colleagues, USFS National Inventory and  
220 Monitoring Applications Center) processed the national plot data to provide our study with  
221 custom products that we employed in this analysis, namely the aboveground live wood biomass  
222 per unit area and its variance for each major forest type, age cohort, productivity class, for each  
223 region shown in Figure 2. We confirmed that the data in this custom delivery were nearly

224 identical to those obtained from other web-based data servers maintained and made available by  
225 the FIA.

226 For this implementation we drive the CASA model with the  $f_{PAR}$  from a smoothed version of  
227 the MODIS MOD15A2 product [Nightingale *et al.*, 2009] for each forest type group as well as  
228 climatological seasonality of monthly weather using NASA Goddard Institute of Space Sciences  
229 (GISS) air temperature anomalies [Hansen *et al.*, 1999] added to a temperature climatology  
230 [Leemans and Cramer, 1991], GISS solar radiation [Zhang *et al.*, 2004], and Global Precipitation  
231 Climatology Project (GPCP) precipitation [Adler *et al.*, 2003]. These meteorological driver data  
232 were sampled at the 1-degree scale while  $f_{PAR}$  was provided at 1 km resolution then averaged for  
233 each forest type within each of the 28 simulation climate domains. As such, we obtain carbon  
234 flux trajectories for each combination of simulation domains ( $n = 28$ ), forest-type group ( $n = 3$  to  
235 10), and productivity class ( $n = 2$ ). Forest type group is specified at a 0.01 degree resolution  
236 obtained from Zhu and Evans [1994] (<http://www.fia.fs.fed.us/library/maps/>). Grid cell-level  
237 fractions of forest land in high and low productivity classes for each forest type and stand age  
238 within each region are specified from county level FIA data.

239 We modified CASA to capture disturbance impacts on the carbon cycle as follows. The post-  
240 disturbance decline and ensuing recovery of  $NPP$  and fractional allocation to wood ( $\tau$ ) are  
241 modeled as:

242

$$243 \quad NPP(t) = NPP_{max}(1 - ce^{-kt}) \quad (4),$$

244

$$245 \quad \tau = \min[1, (t - 1) / 8 \text{ years}] / 3 \quad (5),$$

246

247 where  $t$  is years since disturbance,  $NPP_{max}$  is the climatologically averaged net primary  
248 productivity independent of a disturbance legacy,  $c$  ( $=1.5$ ) determines the magnitude of  
249 disturbance-induced reduction in  $NPP$ ,  $k$  ( $=0.8$ ) determines the rate of  $NPP$  recovery, and  $\min$  is  
250 the minimum operator. We introduced this dynamic recovery of  $NPP$  after disturbance based on  
251 the well documented recovery of  $NPP$  [e.g. *Amiro et al.*, 2000; *Hicke et al.*, 2003]. The dynamics  
252 of allocation were intended to capture initial investment of  $NPP$  into herbaceous biomass with  
253 increasing allocation to woody vegetation with age [e.g. *Jokela et al.*, 2004; *Law et al.*, 2002].

254 In order to parameterize the amount of biomass killed by a disturbance we adopt the following  
255 treatment. Regardless of the pre-disturbance biomass, we set the post-disturbance biomass to  
256 50% of the aboveground live wood biomass reported in the 0-20 year age class. This constrains  
257 early regrowth to pass through the youngest age-class in the FIA sample. We then estimate the  
258 corresponding fraction of live wood, leaves, and roots killed based on the ratio of their abundance  
259 prior to disturbance relative to those immediately after disturbance. Eighty percent of the  
260 disturbance-killed aboveground wood and all of the disturbance-killed leaves are assumed to be  
261 taken off site and entrained into wood products or promptly combusted and are collectively  
262 accounted for as “removals” (fire and harvest), akin to the treatment by *Turner et al.* [1995]. The  
263 remaining 20% of disturbance-killed aboveground wood is subject to on-site post-disturbance  
264 decomposition as it enters the coarse woody debris pool, also consistent with *Turner et al.* [1995].  
265 Disturbance-killed roots decompose on-site, for which 30% of dead coarse roots are assumed to  
266 enter a belowground coarse woody debris pool, and 70% of dead coarse roots and all dead fine  
267 roots enter the soil metabolic and structural pools, broadly consistent with results presented in  
268 [*Gough et al.*, 2007; *Meigs et al.*, 2007]. We note that these and other prescriptions are uncertain,  
269 likely vary among disturbance and forest types, and are the subject of ongoing research. In

summary, biomass killed in a disturbance event is the difference between pre-disturbance biomass and 50% of the 0-20 year biomass reported by the FIA data. Of the killed biomass, 80% of aboveground wood and all leaves are removed (via harvest or fire) and 20% of the killed aboveground wood enters the coarse woody debris pool. The belowground wood and roots killed by disturbance remain on site to decompose. Figure 3 offers an example, in which aboveground biomass is reduced to  $2.5/30 \text{ kg C m}^{-2}$ , or <10%, and 80% of this 90% reduction in biomass is assumed to be removed (harvest or fire) while the other 20% is left to decompose on site.

With this approach it is then possible to estimate biomass removals as:

$$R = A_1 B_{pre} (1 - f_{left}) \quad (6),$$

where  $A_1$  is the area of forested land assigned a stand age of one year based on the FIA age histogram,  $B_{pre}$  is the pre-disturbance aboveground biomass, and  $f_{left}$  (=0.8) is the fraction of biomass left to decompose on-site. Each of these varies by forest type, region, and productivity class. This estimate is subject to errors in the area of forest assigned to this young age class, the age of forests prior to disturbance and correspondingly the biomass pre-disturbance, and uncertainty in the fraction of biomass in disturbed forests that is taken off-site as wood products. Removals from non-stand replacing harvests are not considered in this approach but later in the Discussion section we attempt to quantify the impacts of this assumption.

The next step in our model parameterization involves calculating the wood production – wood age pair that allows the best match to the inventory data of aboveground stock recovery, with the following multi-step procedure. First, we calculate a target aboveground live wood biomass ( $B^*$ , in  $\text{g C m}^{-2}$ ) from the mean in the 100 to 200 year old age classes, including successively younger age classes in 20 year increments to ensure a minimum of two samples. The target age ( $A^*$ , in years) is obtained from the average of old classes sampled to derive  $B^*$ . Second, we approximate

293 the rate of annual aboveground live wood biomass production ( $P_w$ , in  $\text{g C m}^{-2} \text{ a}^{-1}$ ), which is a  
 294 function of  $NPP$  and wood allocation, that would be required to obtain  $B^*$  by  $A^*$  for a range of  
 295 possible wood turnover times ( $A_w$ ) spanning 30 to 300 years in increments of 10 years by solving  
 296 a simplified integral form of the differential equation for biomass with time ( $\text{d}B/\text{d}t = P_w - B/A_w$ ) to  
 297 yield:

298

$$299 \quad P_w = \frac{B^*}{A_w(1 - e^{\frac{-A^*}{A_w}})} \quad (7).$$

300

301 Thus, we obtain an array of possible  $P_w$ - $A_w$  pairs that would grow the target biomass by the target  
 302 age. In a few particular cases this approach yielded implausible wood ages, but with negligible  
 303 consequence for the scales of analyses presented in this study. The third step is to select the pair  
 304 that provides a biomass recovery curve most like the inventory sample assessed as that which  
 305 minimizes the sum of squared error between modeled and sampled aboveground live wood  
 306 biomass. Modeled biomass is calculated at the sample ages ( $t$ , in years) according to:

307

$$308 \quad B(t, A_w) = B_0 e^{\frac{-t}{A_w}} + P_w A_w (1 - e^{\frac{-A^*}{A_w}}) \quad (8),$$

309

310 where  $B_0$  is an assumed initial biomass of  $200 \text{ g C m}^{-2}$ . Lastly, we linearly rescale the model's  
 311 default monthly  $NPP$  values to provide an annual total  $NPP_{max}$  inferred from the fitted rate of  $P_w$ ,  
 312 as:

$$313 \quad NPP_{max} = \frac{P_w}{\tau \alpha} \quad (9),$$

314 where  $\tau$  (=1/3) is the allocation of  $NPP$  to wood and  $\alpha$  (=0.75) is the fraction of this that is  
315 allocated to the aboveground wood pool (stems and branches) instead of belowground (coarse  
316 roots).

317 Following determination of  $P_w$  and  $A_w$  parameters, characteristic carbon flux trajectories  
318 ( $Q_{afp}$ ) are developed from, first, a 1000 year spin-up to steady-state carbon pools. This is followed  
319 by a disturbance prior to the disturbance of interest with 75 years of regrowth for all forest types  
320 except loblolly pine and longleaf / slash pine (30 years) and Douglas-fir (200 years). The age of  
321 trees at harvest is set to be just older than the typical peak in age histograms reported by the FIA  
322 (see Auxiliary 2, Figure A2.3), except where harvest rotations are known to be short (SE and SC  
323 pines), or where harvest over previous decades tended to target old growth forests with high  
324 economic value (Douglas-fir [*Cohen et al.*, 2002]). This ‘pre-disturbance’ is important in that it  
325 establishes the amount of live carbon subject to disturbance-induced disposition, meaning taken  
326 off-site as removals or decomposing on-site. Finally, we simulate the most recent disturbance  
327 after which we allow 200 years of regrowth to characterize carbon dynamics with stand  
328 development. These procedures result in a group of carbon stock age trajectories analogous to  
329 yield tables.

330 We have not modified CASA’s default treatment of heterotrophic respiration emerging from  
331 microbial decomposition of soil and litter carbon and associated transfers among carbon pools.  
332 The general equation for the rate of heterotropic respiration from a specific carbon pool is:

333

334  $Rh_{pool} = C_{pool} k_{pool} W_{resp} T_{resp} M,$  (10),

335

336 where  $C_{pool}$  is the amount of carbon in a pool,  $k_{pool}$  is the pool-specific decay rate constant,  $W_{resp}$   
337 and  $T_{resp}$  control how respiration depends on soil moisture and temperature states, and  $M$  is the  
338 carbon assimilation efficiency of the microbes. Total heterotrophic respiration is the sum of that  
339 from each of the nine detrital pools.

340

### 341 **2.3 Uncertainty Analysis**

342 A formal propagation of uncertainty from sampling errors (coefficient of variation, CV) for  
343 forested area ( $\pm 10$  to 100%) and total aboveground live biomass ( $\pm 10$  to 100%), and volume to  
344 carbon conversion ( $\pm 7\%$ ) are all included. The uncertainty in inventory aboveground live  
345 biomass per unit area is propagated to the predicted fluxes and aboveground live biomass with a  
346 Monte Carlo procedure analogous to Tier 2 uncertainty estimation in the IPCC Good Practice  
347 Guide [IPCC, 2000]. The model was fit to 25 different biomass regrowth trajectories, where each  
348 trajectory was generated from random samples of the normally distributed aboveground live  
349 wood biomass for each age class (25 draws of biomass per unit area from each of 10, 20-year age  
350 classes). Forcing the fitted trajectory to conform to the assumption that biomass increases  
351 monotonically and saturates with age strongly constrains the resultant age-accumulation curves  
352 and their variances (Figure 3). An additional 7% uncertainty is used to account for tree volume to  
353 carbon conversion [Smith and Heath, 2001]. Put together this method involved over 130,000  
354 simulations of age-dependent dynamics of forest carbon fluxes and stocks. The uncertainty of  
355 forest area and aboveground live biomass per unit area is obtained from the FIA data.

356 As shown in Equations 1-3 above, independent uncertainties in the product of flux or stock

357 with area are combined as  $\delta Q_{total} = \left( \frac{\delta Q_{afp}}{Q_{afp}}^2 + \frac{\delta A_{afp}}{A_{afp}}^2 \right)^{1/2}$  [Taylor, 1997]. We adopt a conservative

358 assumption of non-random error propagation for which uncertainty is additive over forest types,

359 productivity classes, and ages, and also additive spatially for a simulation domain, a region, or the  
360 nation. This uncertainty aggregation is analogous to a Tier 1 uncertainty described in the IPCC  
361 Good Practice Guidance [IPCC, 2000].

362 Uncertainty in *NEP* also derives from model structure (not analyzed) as well as model  
363 parameterization of light, moisture, and temperature sensitivity of heterotrophic respiration and/or  
364 *NPP* expressed in the CASA model. As described in Auxiliary Material, Auxiliary 1 Section 1,  
365 we analyzed *NEP* responses to a 2% increase of six representative parameters including the  
366 maximum light use efficiency, moisture dependence of *NPP*, optimal temperature for *NPP*,  
367 turnover time of the slow soil carbon pool, and both the Q10 and moisture dependence of  
368 heterotrophic decomposition of soil carbon. We use a 2% change in parameter value in order to  
369 obtain a detectable response in *NEP* but for ease of discussion the sensitivities are divided by two  
370 and expressed as % change in *NEP* for a 1% change in parameter value (see Auxiliary Material,  
371 Auxiliary 1).

372

### 373 **3. Results**

#### 374 **3.1 Carbon Trajectories**

375 Using CASA as a controlled growth model accurately reproduces the accumulation of  
376 aboveground forest carbon stocks with time since a stand replacing disturbance as informed by  
377 FIA data (Figure 3), imposing a powerful, albeit partial, observational constraint on net ecosystem  
378 carbon flux trajectories with stand age. Additional data on litter, woody debris and soil carbon  
379 dynamics would provide much needed additional constraints on estimated ecosystem C dynamics.

380 More rapid regrowth of aboveground stocks in the high productivity class causes higher  
381 amplitude trajectories for carbon stocks and fluxes (Figure A2.1, A2.2, Auxiliary 2) with larger

382 post-disturbance sources that give way to stronger sinks with ensuing forest regrowth. The Monte  
383 Carlo simulation approach provides an envelope of trajectories (Figure 3) that enables formal  
384 uncertainty propagation through all scales of the analysis (regional forest types to conterminous  
385 U.S. forestlands). Absolute uncertainty surrounding *NEP* tends to peak where forest uptake is  
386 maximum (peak *NEP*) and then diminishes with forest age (Figure 3). An important exception,  
387 not shown in Figure 3, is the often large uncertainty in carbon emission in the years immediately  
388 following disturbance; large because of variation in the pre-disturbance carbon stocks and the  
389 amount of dead wood that decomposes on-site. The timing of *NEP* crossover from source to sink  
390 is surprisingly insensitive to variability in biomass accumulation (not shown), and generally  
391 occurs at ages <20 years (e.g. Figure 3 and Figure A2.1, A2.2, Auxiliary 2) consistent with many  
392 reported chronosequence fluxes [e.g. *Bond-Lamberty et al.*, 2004; *Gough et al.*, 2007; *Goulden et*  
393 *al.*, 2011; *Law et al.*, 2004; *Litvak et al.*, 2003; *Noormets et al.*, 2007; *Pregitzer and Euskirchen*,  
394 2004]. Patterns of post-disturbance uptake of carbon in regrowing forests vary widely across  
395 regions of the conterminous U.S. as well as by forest type group and productivity class (Figure  
396 A2.1, A2.2, Auxiliary 2). Forest inventory data describing the recovery of aboveground live wood  
397 biomass carbon with stand development act as a strong constraint on the modeled carbon cycle  
398 including the rates of litter and soil carbon turnover and decay.

399 Our analysis of the sensitivity of the model to parameters revealed that nearly all of the  
400 sensitivities are less than 1% indicating general dampening of parameter perturbations and  
401 suggesting that uncertainties in these parameterizations do not expand as they propagate through  
402 to modeled *NEP* (see Auxiliary 1, Table A1.1). Model structure and parameter uncertainties are  
403 not included in our analysis but are expected to add about  $\pm 10\%$  based partly on a sensitivity  
404 analysis presented in Auxiliary Material 1.

405

406 **3.2 Continental Patterns**

407     Regional variations in disturbance rates and *NEP* across the conterminous U.S. reflect  
408     harvesting practices and regional climates (Tables 1 and Table A2.1, Figures A2.1, A2.2, A2.3).  
409     Forests growing in relatively dry settings (e.g. Rocky Mountain South (RMS)) have low *NEP*,  
410     contrasted by high carbon sequestration rates in the Pacific Southwest and Northwest, as well as  
411     Southeastern and South Central regions (Table A2.1). The largest rates of disturbance, and the  
412     largest sinks of carbon stimulated by forest recovery from recent disturbance (“regrowth sinks”),  
413     are in Southeastern (SE), South Central (SC), and Pacific Northwest (PNW) regions. These  
414     regional biologically driven sinks do not reflect net biome productivity because recovery  
415     trajectories do not include the fate of disturbance-induced carbon removals such as carbon taken  
416     offsite to lumber, pulp and paper mills or released promptly on-site by natural and anthropogenic  
417     fires (see schematic in Figure 1). This is addressed further in the discussion where we present the  
418     forest-to-atmosphere carbon exchange.

419     At the continental scale, the biological recovery sink (*NEP*) is estimated to be  $164 \pm 28 \text{ Tg C a}^{-1}$   
420     (Table 1), or about  $71 \text{ g C m}^{-2} \text{ a}^{-1}$  averaged for the 230 million hectares of forestland represented  
421     here. Nearly all (84%) of this *NEP* sink results from net growth of live carbon stocks with only a  
422     small fraction shared among soil carbon (6%), litter carbon (2%), and coarse woody debris (8%)  
423     stocks (Table 2). Our sample includes 93% of the conterminous U.S. forestland, reported to be  
424     250 million hectares [EPA, 2008]. Our analysis did not include the Northern Prairie States region  
425     (~6% of total area and ~5% of total carbon) because the effort was originally connected to a  
426     Landsat remote sensing analysis whose random sample did not draw Landsat scenes for this  
427     region. As verification, our stand-age histograms by region generally correspond well with a

428 similar presentation of the same basic data as recently published by *Pan et al.* [2011]. Comparing  
429 to regional statistics of forest area and live biomass reported in *EPA* [2008] we find good  
430 correspondence overall (Table 1).

431 The estimated uncertainty arising from forest area, aboveground wood biomass, and  
432 conversion of diameter measurements to volume and carbon produced relatively small uncertainty  
433 estimates in our biomass and fluxes. This is partly due to the continuous, monotonically  
434 increasing, and saturating growth form imposed by the process-model approach. This functional  
435 form is more plausible than one that would allow abrupt increases and decreases in aboveground  
436 live wood biomass with stand development (i.e. stand age) as are commonly found in the  
437 inventory data when arrayed as a chronosequence (e.g. Figure 3, 110- 150 year biomass).  
438 Imposing the model's growth form has the effect of filtering out some of the variance inherent in  
439 chronosequence trajectories of biomass with stand age. Other uncertainties arising from model  
440 structure and assumptions about disturbance severity/type, age, partial cutting, natural wood  
441 turnover, and a possible age-related decline in productivity are evaluated by judging the impacts  
442 of these factors on model output through sensitivity analyses (see Auxiliary 1, Section 2).

443 We used the 1km forest type map to produce a gridded map of *NEP* and its uncertainty (from  
444 variances in FIA data) for the conterminous US (Figure 4). Within each region each forest type  
445 considered was assigned the regional estimate of *NEP* for that forest type and region. Regional  
446 forest *NEP* sinks range from  $>25$  to  $200 \text{ g C m}^{-2} \text{ a}^{-1}$  with eastern and western forests generally  
447 ranging from  $75$  to  $100 \text{ g C m}^{-2} \text{ a}^{-1}$ . The RMS region is predicted to be uniformly  $<50 \text{ g C m}^{-2} \text{ a}^{-1}$ .  
448 The discontinuities conforming to state borders between West Virginia and Virginia and between  
449 Washington and Idaho occur because the same forest types in each neighboring region have  
450 regionally specific and different growth and disturbance rates.

451 As an independent evaluation of our predicted stocks and fluxes we compared our results with  
452 five available studies on chronosequences for forest types in the conterminous US. These studies  
453 sometimes do not include estimates of both fluxes and stocks for different aged forests and  
454 estimates used various biometric and flux measurement approaches. The small number of sites  
455 with available data, variability in the data, and issues of extrapolating fine scale measurements to  
456 regional responses do not justify quantitative comparisons and demonstrate the need for more of  
457 these types of measurements and for finer scale modeling. The results of these comparisons are  
458 shown in Auxiliary 3, Figure A3.1. Agreement varies widely between the comparisons at the  
459 different sites/regions.

460

## 461 **4. Discussion**

462 Comparing estimates of the conterminous U.S. forest *NEP* sink from multiple studies (Table  
463 3) reveals a general separation between age-accumulation and stock-change methods. This  
464 comparison spans estimates for the 1980s to more recent years (e.g. 2005-2006), but this may be  
465 justified because atmospheric inversions seem to indicate a long term mean sink in North  
466 America during the '80's and '90's but with large interannual variability [Baker *et al.*, 2006].  
467 Four of the six age dependent analyses that seek to represent carbon emissions and sequestration  
468 with post-disturbance recovery provide lower estimates of the forest *NEP* sink when compared to  
469 the four stock-change analyses, with  $82 \text{ g C m}^{-2} \text{ a}^{-1}$  versus  $154 \text{ g C m}^{-2} \text{ a}^{-1}$  averaged across their  
470 respective studies, or  $189 \text{ Tg C a}^{-1}$  versus  $354 \text{ Tg C a}^{-1}$  when integrated across US forest area.  
471 This is even true when process-oriented studies rely on forest inventory data to prescribe the rate  
472 of aboveground carbon stock recovery with time, as well as the area of forest of different ages.

473 For example, regarding *NEP* alone we find general agreement with *Turner et al.* [1995] who  
474 reported 203 Tg C a<sup>-1</sup> compared to our estimate of 164 Tg C a<sup>-1</sup>. In contrast, the *EPA* [2008]  
475 stock-change estimate of forest *NEP* is twice as large as this study's age-accumulation result (335  
476 compared to 164 Tg C a<sup>-1</sup>, Table 2). The disparity between the stock-change method and these  
477 other, age-accumulation results is likely due to large annual to decadal increases in stocks  
478 measured in the inventory that then implies greater *NEP* (regrowth). What causes this general  
479 disagreement remains unclear, though growth enhancement is a plausible explanation of the  
480 difference, consistent with recent publications [*Cole et al.*, 2010; *Luyssaert et al.*, 2010;  
481 *McMahon et al.*, 2010; *Thomas et al.*, 2009]. Effects of growth enhancement are implicit in the  
482 stock-change method but not well incorporated in the age-accumulation methods that emphasize  
483 effects of regrowth dynamics, even when these methods rely on inventory-derived  
484 chronosequences to constrain biomass accumulation as in the present study (see Auxiliary 4 for  
485 an illustration of this). There is also one study reported in Table 3 including only the effects of  
486 climate and CO<sub>2</sub> fertilization based on an ensemble of models for the conterminous US [*Schimel*  
487 *et al.*, 2000]. If this sink were added to the forest recovery (age-yield table) estimates the results  
488 would be more in line with the stock change approach.

489 We note that the *EPA* [2008] estimate of total removals is 38% higher than that estimated with  
490 our modeling approach (=162/117, Table 3). About half of the difference is due to elevated fire  
491 emissions reported in *EPA* [2008], however this estimate is much higher than the rate of forest  
492 fire emissions being reported elsewhere (e.g. *van der Werf et al.*[2010]). This difference  
493 translates directly into the *NEP* estimated from the stock change method, and elevates the *EPA*  
494 [2008] estimate by 20 TgC a<sup>-1</sup> relative to the estimate from our approach. The *EPA* [2008] report  
495 also estimates 25 TgC a<sup>-1</sup> greater removals by harvest. There are two ways we could adjust our

496 methodology to try to match this rate of removal. We could either, a) increase the amount of  
497 biomass removed by disturbances on average by increasing the age and hence biomass of  
498 disturbed forests, b) increase the amount of biomass removed on average by removing a larger  
499 fraction of pre-disturbance biomass and leaving less to decompose on site, or c) increase the area  
500 of forests disturbed by increasing the young-aged fraction of forests if we believe the stand age  
501 attribute offers a biased representation. All of these would increase removals but they would have  
502 different effects on *NEP*. The first option would decrease *NEP* because more disturbance-killed  
503 material would be left on-site to decompose and be emitted from forests. The second approach  
504 would increase *NEP* because of reduced on-site decomposition. The third approach would  
505 decrease *NEP* because a larger fraction of forested area would be concentrated at young stand  
506 ages (<15 year old) where *NEP* is either a large negative value or near zero (Figures A2.1, A2.2).  
507 And in the extreme case that we simply adjusted our *NEP* estimate upwards to cover the  
508 difference in removals, the *EPA* [2008] estimate would still be 126 TgC a<sup>-1</sup> higher than the  
509 estimate emerging from our age-accumulation method.

510 A term-by-term comparison between stock changes reported from inventory methods and  
511 those derived in the current study's age-accumulation approach indicates that a change in live  
512 carbon stocks makes up a large portion of the difference in *NEP* estimated with the two methods  
513 (Table 2). Annual increases in soil carbon, coarse woody debris, and litter pools are also  
514 noticeably lower in the present analysis compared to those reported by the *EPA* [2008; 2010]  
515 (Table 3). Because our method, necessarily, produces aboveground live wood biomass and forest  
516 area estimates that are consistent with, or wholly derived from, the inventory itself (Table 1), our  
517 relatively low estimate of annual changes in live stocks (Table 2) does not appear to be caused by  
518 underestimation of a) stocks, or b) forest area. These differences translate to the full forest sector-

519 atmosphere net exchange, whereby the stock-change method estimates a much larger forest sector  
520 C sink than obtained with this study's age-accumulation approach (Table 2).

521 Our maps of conterminous US forest *NEP* and its uncertainty (Figure 4) are one of the first of  
522 which we are aware (though see *Woodbury et al.* [2007]) and will be used in further study of the  
523 impact of the forest disturbance fluxes on atmospheric CO<sub>2</sub> as a boundary flux for atmospheric  
524 transport models much as gridded fire, fossil fuel burning, and ocean CO<sub>2</sub> fluxes are prescribed in  
525 forward and inverse atmospheric modeling [e.g. *Peters et al.*, 2007]. Complete accounting of  
526 forest sector fluxes would additionally require maps of fire [e.g. *van der Werf et al.*, 2010] and  
527 wood products emissions. These studies will allow assessment of the detection limits for the  
528 magnitude and spatial variability of sinks in top-down studies.

529 This study's approach imposed a number of simplifying assumptions that were necessary  
530 given the initial scope of our work. Below we address some of these and their potential  
531 implications regarding interpretation of our results.

532 • We assume characteristic regrowth trajectories regardless of disturbance type even though  
533 the nature of post-disturbance carbon dynamics is sure to vary between fire, harvest,  
534 hurricane, and the severity of disturbance. For instance, around twice as much coarse  
535 woody debris (CWD) may remain on site after a severe fire compared to clear-cut harvest  
536 [*Tinker and Knight*, 2000]. This remaining detritus provides a source of CO<sub>2</sub> for a  
537 prolonged period after disturbance. Using data reported by *Smith et al.* [2009] and the  
538 National Interagency Fire Center (to account for Alaskan fires) we estimate that for the year  
539 2004 the ratio of burned area to harvested area in the eastern US was about 0.30 compared  
540 to 0.46 in the west. In terms of carbon removals though, our forest fire estimates from the  
541 Global Fire Emissions Database v3 (10 Tg C a<sup>-1</sup>) are much smaller than our estimated

542 harvest removals ( $107 \text{ Tg C a}^{-1}$ ). Because the total removals are dominated by harvest, as is  
543 the total area disturbed, accounting for differences caused by fire versus harvest would not  
544 significantly change our results or conclusions. Furthermore, some but not all of this  
545 variation is captured by the Monte Carlo approach, as well as with stratification by site  
546 productivity and across regions. Partial disturbances such as defoliation events are not  
547 represented with the current methodology, and discussed further below.

- 548 • Our assumption of equivalence between forest age and time since disturbance does not  
549 account for the effects of partial disturbance that allows older aged trees to remain among  
550 regenerating cohorts or the dynamic state of old forests that have reached the age of natural  
551 mortality and reestablishment. This particular issue has been examined by *Bradford et al.*  
552 [2008] for a subalpine forest system. In that study a large part of the age versus years since  
553 disturbance discrepancy arose in stands undisturbed for long periods of time (>200 years),  
554 longer than what we analyze in this work. From FIA data we estimate that about 3% of  
555 forested land is >200 years old for conterminous US.
- 556 • Our analysis is sensitive to biases in the ages associated with the aboveground live wood  
557 biomass trajectories, as explored in an extensive sensitivity analysis described in Auxiliary  
558 1, Section 2. For instance, if the FIA ages are older (younger) than actual stand ages, our  
559 predicted recovery sink is underestimated (overestimated). This, of course, is an issue with  
560 any approach proposing to use FIA age structure information to estimate fluxes and stocks  
561 [e.g. *Pan et al.*, 2011]. Despite this sensitivity, we note that bias in stand age is not likely to  
562 be large enough to explain the major differences between the stock-change and age-  
563 accumulation methods (Table A1.2).

- 564 • The FIA data we used to construct aboveground live wood biomass trajectories include the  
565 effects of partial cuts, which are a significant component of disturbance in US forests  
566 contributing >50% of the total harvested area [Smith *et al.*, 2009]. Reported stand ages  
567 reflect the trees not cut while the plot level biomass will be lower in these cases producing  
568 lower regional aboveground live wood biomass for mid and older aged stands. These partial  
569 cutting practices (e.g. salvage logging, selective logging, thinning), which remove biomass  
570 from forested plots without resetting the FIA-recorded stand age, could have a substantial  
571 influence on the forest *NEP* estimate. The implicit inclusion of plots that experienced  
572 partial cutting (not fully stocked) likely results in correct biomass estimates but lowers the  
573 slope of regrowth trajectories resulting in some underestimation of *NEP*. In an extensive  
574 sensitivity analysis (Auxiliary 1, Section 2) we find strong sensitivity to such biases, with a  
575 10% elevation of biomass leading to a 14% elevation of conterminous US forest *NEP*. This  
576 is equivalent to a  $2.3 \text{ Tg C a}^{-1}$  increase in *NEP* for each 1% increase in biomass. Despite  
577 this large sensitivity to biomass trajectories, to account for the approximately  $160 \text{ Tg C a}^{-1}$   
578 difference, the reported biomass would need to have been underestimated by 70% ( $=160 \text{ Tg}$   
579  $\text{C a}^{-1} / 2.3 \text{ Tg C a}^{-1}$  per 1% increase in biomass). Additional sensitivity analyses examining  
580 effects of natural, partial disturbances that lead to wood turnover and on site decomposition  
581 (e.g. ice storms, blowdowns, insect damage) indicate that they are also unlikely to present a  
582 large error/bias in our estimate.
- 583 • We do not take into account annual changes in forest area which could contribute to the  
584 discrepancy between recovery and stock change approaches. The *EPA* [2008] reports  
585 indicate that forest area has been increasing at a rate of  $0.24\% \text{ a}^{-1}$  since 1990. If we assume  
586 that new forests would range between 1 to  $5 \text{ kg C m}^{-2}$  over an age range of 0 to 20 years

587 (e.g. see Figure 3) then the average accumulation rate for these forest would be about 250 g  
588 C m<sup>-2</sup> a<sup>-1</sup>. Correcting this for the increase in forest area produces an added 1.7 Tg C a<sup>-1</sup> sink,  
589 indistinguishable within the uncertainties of our method.

590 • It has been proposed that forest carbon sinks may be driven by long term trends in  
591 temperature, precipitation, nitrogen deposition, and atmospheric CO<sub>2</sub>. Responses to these  
592 trends are embedded in the biomass-age trajectories from the inventories in complex ways  
593 and more recent increases in growth may not be accounted for in our approach (see  
594 Auxiliary 4 for a thorough examination of this). Others have addressed this and concluded  
595 that forests are not responding in a systematic way to these trends [Caspersen *et al.*, 2000],  
596 that forest inventory data are not precise enough to resolve expected responses to trends  
597 [Joos *et al.*, 2002], and that a smaller number of inventory measurements on forests of  
598 known disturbance history do indeed show strong trends in growth enhancement correlated  
599 with trends in temperature and atmospheric CO<sub>2</sub> [McMahon *et al.*, 2010; Thomas *et al.*,  
600 2009]. In a study of global terrestrial carbon sinks using CASA, Thompson *et al.* [1996]  
601 showed that in order to obtain a terrestrial carbon sink of ~2 Pg C/yr broadly consistent with  
602 top-down sink estimates, *NPP* has to undergo a sustained increase of 0.18% per annum.  
603 Similar estimates have been reported by others [e.g. Joos *et al.*, 2002]. Our own sensitivity  
604 analysis (not shown) showed that a sustained increase in *NPP* of 0.2% per annum would  
605 increase live biomass in a typical 60 year old forest by approximately 5% and is thus a weak  
606 or undetectable signal in a biomass chronosequence. A 0.2% annual increase in *NPP* is  
607 implausibly large sensitivity of photosynthesis to CO<sub>2</sub> ( $dNPP/NPP \times CO_2/dCO_2$  of ~0.96, or  
608 near proportional response) and would require other positive feedback mechanisms such as  
609 nitrogen fertilization and/or climate trends to operate in parallel. We conclude that plausible

610 responses of forest sinks to climate and CO<sub>2</sub> or N cannot be resolved with FIA biomass-age  
611 trajectories alone such as those we utilize here and that have been proposed by others [e.g.  
612 *Pan et al.*, 2011].

613 The approach described here is also sensitive to uncertain parameters including rates of wood  
614 mortality and coarse woody debris decomposition, as well as the amount of dead aboveground  
615 and belowground biomass left to decompose onsite following disturbance. It lacks a standing  
616 dead wood pool that may be important because it decomposes much more slowly than dead wood  
617 in contact with the forest floor [e.g. *Harmon and Hua*, 1991; *Harmon et al.*, 2004; *Janisch et al.*,  
618 2005]. In our ongoing efforts, literature is being exhaustively explored to better constrain these  
619 and other parameters and processes. Additional effort is being invested in attributing disturbances  
620 to particular drivers based on spatial and geospatial records of fire and insect outbreaks. While  
621 valuable, it is unlikely that such refinements and constraints will reconcile the large differences  
622 between the age-accumulation and stock-change approaches, something that may benefit from a  
623 close collaboration with inventory experts to clarify differences of approach and accounting, as  
624 well as more comprehensive assessment of possible growth enhancement effects. Future efforts  
625 at improving this study's approach will include more detailed prescriptions of type and severity of  
626 disturbances, further comparisons with site observations as they become available, and analyses  
627 of top-down atmospheric constraints on source/sink magnitude and distributions. Estimates  
628 would also be better constrained if additional data on litter, dead wood and soil organic carbon  
629 dynamics were available from field studies.

## 630 **5. Conclusions**

631 Forest Inventory and Analysis data provide unique and valuable information about  
632 disturbance history and associated carbon stocks and fluxes with forest recovery. By using these

633 data to constrain forest growth rates in a carbon cycle model, this study provides a more detailed  
634 estimate of carbon sources and sinks from recent forest disturbance and recovery across regions  
635 and forest types of the US. One of our key findings is a much smaller net sink of carbon in  
636 conterminous US forests than previously estimated with the stock-change approach as used in  
637 UNFCCC reporting [EPA, 2008]. The source of across study inconsistencies among national  
638 estimates of stocks and fluxes remains largely unexplained. The paucity of observed net  
639 ecosystem productivity and biomass chronosequences limits our ability to evaluate modeled  
640 responses. These types of observations are critically needed in order to adequately test models  
641 representing disturbance and subsequent recovery.

642

643 **Acknowledgements:** We thank Charles (Chip) Scott and his team at the USFS National  
644 Inventory and Monitoring Applications Center for providing us with FIA data. We acknowledge  
645 helpful discussions and/or comments from Skee Houghton, Jim Randerson, Warren Cohen, Linda  
646 Heath, and an anonymous reviewer. This work was funded by NASA NNH05ZDA001N, North  
647 American Carbon Program. In addition, CAW was supported by the US National Science  
648 Foundation under grant ATM-0910766.

649 **References**

- 650 Adler, R., et al. (2003), The Version-2 Global Precipitation Climatology Project (GPCP) monthly  
651 precipitation analysis (1979-present), *Journal of Hydrometeorology*, 4(6), 1147-1167.
- 652 Amiro, B. D., et al. (2000), Net primary productivity following forest fire for Canadian  
653 ecoregions, *Canadian Journal of Forest Research-Revue Canadienne De Recherche  
654 Forestiere*, 30(6), 939–947.
- 655 Amiro, B. D., et al. (2011), Ecosystem carbon dioxide fluxes after disturbance in forests of North  
656 America, *Journal of Geophysical Research*, 115, G00K02.
- 657 Baker, D. F., et al. (2006), TransCom 3 inversion intercomparison: Impact of transport model  
658 errors on the interannual variability of regional CO<sub>2</sub> fluxes, 1988-2003, *Global  
659 Biogeochemical Cycles*, 20, GB1002, doi:1010.1029/2004GB002439.
- 660 Barford, C., et al. (2001), Factors controlling long-and short-term sequestration of atmospheric  
661 CO<sub>2</sub> in a mid-latitude forest, *Science*, 294, 1688-1691.
- 662 Bechtold, W. A., and P. L. Patterson (2005), The enhanced Forest Inventory and Analysis  
663 program—national sampling design and estimation procedures, SRS GTR-80, USDA Forest  
664 Service, Southern Research Station, Asheville, North Carolina, USA.
- 665 Birdsey, R. A., and L. S. Heath (1995), Carbon changes in U.S. forests, in *Productivity of  
666 America's Forests and Climate Change*, edited by L. A. Joyce, Fort Collins.
- 667 Bond-Lamberty, B., et al. (2004), Net primary production and net ecosystem production of a  
668 boreal black spruce wildfire chronosequence, *Global Change Biology*, 10, 473-487.
- 669 Bousquet, P., et al. (2000), Regional changes in carbon dioxide fluxes of land and oceans since  
670 1980, *Science*, 290(5495), 1342-1346.

- 671 Bradford, J. B., et al. (2008), Tree age, disturbance history, and carbon stocks and fluxes in  
672 subalpine Rocky Mountain forests, *Global Change Biology*, 14, 2882-2897.
- 673 Caspersen, J. P., et al. (2000), Contributions of land-use history to carbon accumulation in US  
674 forests, *Science*, 290(5494), 1148-1151.
- 675 Chapin III, F. S., et al. (2006), Reconciling carbon-cycle concepts, terminology, and methods,  
676 *Ecosystems*, 9, 1041-1050.
- 677 Cohen, W., et al. (2002), Characterizing 23 years (1972-95) of stand replacement disturbance in  
678 western Oregon forests with Landsat imagery, *Ecosystems*(122-137).
- 679 Cole, C. T., et al. (2010), Rising concentrations of atmospheric CO<sub>2</sub> have increased growth in  
680 natural stands of quaking aspen (*Populus tremuloides*), *Global Change Biology*, 16, 2186-  
681 2197.
- 682 EPA (2008), Inventory of U.S. Greenhouse Gas Emissons and Sinks: 1990-2006, United States  
683 Environmental Protection Agency, edited, Washington, DC.
- 684 EPA (2010), Inventory of U.S. Greenhouse Gas Emissons and Sinks: 1990-2008, United States  
685 Environmental Protection Agency, Washington, DC.
- 686 Fan, S., et al. (1998), A large terrestrial carbon sink in North America implied by atmospheric and  
687 oceanic carbon dioxide data and models, *Science*, 282, 442-446.
- 688 Gough, C. M., et al. (2007), The legacy of harvest and fire on ecosystem carbon storage in a  
689 northern temperate forest, *Global Change Biology*, 13, 1935-1949.
- 690 Goulden, M. L., et al. (2011), Patterns of NPP, GPP, respiration, and NEP during boreal forest  
691 succession, *Global Change Biology*, 17, 855-871.
- 692 Goward, S. N., et al. (2008), Forest disturbance and North American carbon flux, *EOS*  
693 *Transactions, American Geophysical Union*, 89(11), 105-106.

- 694 Gurney, K. R., et al. (2002), Towards robust regional estimates of CO<sub>2</sub> sources and sinks using  
695 atmospheric transport models, *Nature*, 415(6872), 626-630.
- 696 Hansen, J., et al. (1999), GISS analysis of surface temperature change, *Journal of Geophysical*  
697 *Research*, 104, 30997-31022.
- 698 Harmon, M. E., and C. Hua (1991), Coarse woody debris dynamics in two old-growth  
699 ecosystems, *BioScience*, 41(9), 604-610.
- 700 Harmon, M. E., et al. (2004), Production, respiration, and overall carbon balance in an old-growth  
701 Pseudotsuga-Tsuga forest ecosystem, *Ecosystems*, 7, 498-512.
- 702 Hicke, J. A., et al. (2003), Postfire response of North American boreal forest net primary  
703 productivity analyzed with satellite observations, *Global Change Biology*, 9(8), 1145-1157.
- 704 Houghton, R. A. (1999), The annual net flux of carbon to the atmosphere from changes in land  
705 use 1850-1990, *Tellus Series B-Chemical and Physical Meteorology*, 51(2), 298-313.
- 706 Houghton, R. A., et al. (1999), The US carbon budget: Contributions from land-use change,  
707 *Science*, 285(5427), 574-578.
- 708 Houghton, R. A. (2003), Why are estimates of the terrestrial carbon balance so different?, *Global*  
709 *Change Biology*, 9(4), 500-509.
- 710 Hurtt, G. C., et al. (2002), Projecting the future of the US carbon sink, *Proceedings of the*  
711 *National Academy of Sciences of the United States of America*, 99(3), 1389-1394.
- 712 IPCC (2000), Intergovernmental Panel on Climate Change Good Practice Guidance and  
713 Uncertainty Management in National Greenhouse Gas Inventories, Institute for Global  
714 Environmental Strategies, Japan.
- 715 Ito, A., et al. (2008), Can we reconcile differences in estimates of carbon fluxes from land-use  
716 change and forestry for the 1990s?, *Atmos. Chem. Phys.*, 8, 3291-3310.

- 717 Janisch, J. E., et al. (2005), Decomposition of coarse woody debris originating by clearcutting of  
718 an old-growth conifer forest, *Ecoscience*, 12(2), 151-160.
- 719 Jokela, E. J., et al. (2004), Production dynamics of intensively managed loblolly pine stands in the  
720 southern United States: a synthesis of seven long-term experiments, *Forest Ecology and*  
721 *Management*, 192, 117-130.
- 722 Joos, F., et al. (2002), Growth enhancement due to global atmospheric change as predicted by  
723 terrestrial ecosystem models: consistent with US forest inventory data, *Global Change*  
724 *Biology*, 8, 299-303.
- 725 Kaminski, T., et al. (2001), Inverse modeling of atmospheric carbon dioxide fluxes, *Science*, 294,  
726 259.
- 727 King, A. W., et al. (2007), The First State of the Carbon Cycle Report (SOCCR): The North  
728 American Carbon Budget and Implications for the Global Carbon Cycle, *NOAA, National*  
729 *Climatic Data Center, Asheville, NC*.
- 730 Koerner, C. (2003), Slow in, rapid out - carbon flux studies and Kyoto targets, *Science*, 300,  
731 1242-1243.
- 732 Law, B. E., et al. (2002), Environmental controls over carbon dioxide and water vapor exchange  
733 of terrestrial vegetation, *Agricultural and Forest Meteorology*, 113(1-4), 97-120.
- 734 Law, B. E., et al. (2003), Changes in carbon storage and fluxes in a chronosequence of ponderosa  
735 pine, *Global Change Biology*, 9, 510-524.
- 736 Law, B. E., et al. (2004), Disturbance and climate effects on carbon stocks and fluxes across  
737 Western Oregon USA, *Global Change Biology*, 10, 1429-1444.

- 738 Leemans, R., and W. P. Cramer (1991), The IIASA Database for Mean Monthly Values of  
739 Temperature, Precipitation and Cloudiness of a Global Terrestrial Grid, RR-91-18, 62 pp,  
740 IIASA, Laxenburg, Austria.
- 741 Litvak, M., et al. (2003), Effect of stand age on whole ecosystem CO<sub>2</sub> exchange in the Canadian  
742 boreal forest, *Journal of Geophysical Research*, 108(D3), 8225.
- 743 Luysaert, S., et al. (2010), The European carbon balance. Part 3: forests, *Global Change Biology*,  
744 16, 1429-1450.
- 745 McGuire, A. D., et al. (2001), Carbon balance of the terrestrial biosphere in the twentieth century:  
746 Analyses of CO<sub>2</sub>, climate and land use effects with four process-based ecosystem models,  
747 *Global Biogeochemical Cycles*, 15(1), 183-206.
- 748 McMahon, S. M., et al. (2010), Evidence for a recent increase in forest growth, *Proceedings of  
749 the National Academy of Sciences*, 107(8), 3611-3615.
- 750 Meigs, G. W., et al. (2007), Forest fire impacts on carbon uptake, storage, and emission: The role  
751 of burn severity in the Eastern Cascades, Oregon, *Ecosystems*, 12, 1246-1267.
- 752 Myneni, R. B., et al. (2001), A large carbon sink in the woody biomass of Northern forests,  
753 *Proceedings of the National Academy of Sciences of the United States of America*, 98(26),  
754 14784-14789.
- 755 Nemani, R., et al. (2002), Recent trends in hydrologic balance have enhanced the terrestrial  
756 carbon sink in the United States, *Geophysical Research Letters*, 29(10).
- 757 Nightingale, J. M., et al. (2009), Temporally smoothed and gap-filled MODIS land products for  
758 carbon modelling: application of the fPAR product, *International Journal of Remote Sensing*,  
759 30, 1083-1090.

- 760 Noormets, A., et al. (2007), Age-related changes in forest carbon fluxes in a managed northern  
761 Wisconsin landscape, *Ecosystems*, 10, 187-203.
- 762 Odum, E. (1969), The strategy of ecosystem development, *Science*, 164, 262-270.
- 763 Pacala, S. W., et al. (2001), Consistent land- and atmosphere-based US carbon sink estimates,  
764 *Science*, 292(5525), 2316-2320.
- 765 Pan, Y., et al. (2009), Separating effects of changes in atmospheric composition, climate and  
766 land-use on carbon sequestration of U.S. Mid-Atlantic temperate forests, *Forest Ecology and  
767 Management*, 259, 151-164.
- 768 Pan, Y., et al. (2011), Age structure and disturbance legacy of North American forests,  
769 *Biogeosciences*, 8, 715-732.
- 770 Peters, W., et al. (2007), An atmospheric perspective on North American carbon dioxide  
771 exchange: CarbonTracker, *Proceedings of the National Academy of Sciences of the United  
772 States of America*, 104, 18925-18930.
- 773 Potter, C. S., et al. (1993), Terrestrial Ecosystem Production - a Process Model-Based on Global  
774 Satellite and Surface Data, *Global Biogeochemical Cycles*, 7(4), 811-841.
- 775 Pregitzer, K. S., and E. S. Euskirchen (2004), Carbon cycling and storage in world forests: Biome  
776 patterns related to forest age, *Global Change Biology*, 10, 2052-2077.
- 777 Randerson, J. T., et al. (1996), Substrate limitations for heterotrophs: Implications for models that  
778 estimate the seasonal cycle of atmospheric CO<sub>2</sub>, *Global Biogeochemical Cycles*, 10(4), 585-  
779 602.
- 780 Richter, D. D., et al. (1999), Rapid accumulation and turnover of soil carbon in a re-establishing  
781 forest, *Nature*, 400, 56-58.

- 782 Schimel, D., et al. (2000), Contribution of increasing CO<sub>2</sub> and climate to carbon storage by  
783 ecosystems in the United States, *Science*, 287(5460), 2004-2006.
- 784 Schimel, D. S., et al. (2001), Recent patterns and mechanisms of carbon exchange by terrestrial  
785 ecosystems, *Nature*, 414(6860), 169-172.
- 786 Schwalm, C. R., et al. (2007), A method for deriving net primary productivity and component  
787 respiratory fluxes from tower-based eddy covariance data: a case study using a 17-year data  
788 record from a Douglas-fir chronosequence, *Global Change Biology*, 13, 370-385.
- 789 Smith, J. E., and L. S. Heath (2001), Identifying influences on model uncertainty: An application  
790 using a forest carbon budget model, *Environmental Management*, 27(2), 253-267.
- 791 Smith, W. B., et al. (2009), Forest Resources of the United States, *Gen. Tech. Rep. WO-78,*  
792 *Washington, DC: USDA, Forest Service, Washington Office.*, 336p.
- 793 Stephens, B. B., et al. (2007), Weak northern and strong tropical land carbon uptake from vertical  
794 profiles of atmospheric CO<sub>2</sub>, *Science*, 316, 1732-1735.
- 795 Tans, P. P., et al. (1990), Observational Constraints on the Global Atmospheric Co<sub>2</sub> Budget,  
796 *Science*, 247(4949), 1431-1438.
- 797 Taylor, J. R. (1997), *An introduction to error analysis: The study of uncertainties in physical*  
798 *measurements 2nd ed.*, 326 pp., University Science Books, Sausalito, California.
- 799 Thomas, R. Q., et al. (2009), Increased tree carbon storage in response to nitrogen deposition in  
800 the US, *Nature-Geoscience*, 3, 13-17.
- 801 Thompson, M. V., et al. (1996), Change in net primary production and heterotrophic respiration:  
802 How much is necessary to sustain the terrestrial carbon sink?, *Global Biogeochemical Cycles*,  
803 10(4), 711-726.

- 804 Thornton, P. E., et al. (2002), Modeling and measuring the effects of disturbance history and  
805 climate on carbon and water budgets in evergreen needleleaf forests, *Agricultural and Forest*  
806 *Meteorology*, 113(1-4), 185-222.
- 807 Tinker, D. B., and D. H. Knight (2000), Coarse woody debris following fire and logging in  
808 Wyoming lodgepole pine forests, *Ecosystems*, 3, 472-483.
- 809 Turner, D. P., et al. (1995), A carbon budget for forests of the conterminous United States,  
810 *Ecological Applications*, 5(2), 421-436.
- 811 van der Werf, G. R., et al. (2010), Global fire emissions and the contribution of deforestation,  
812 savanna, forest, agricultural, and peat fires (1997-2009). *Atmos. Chem. Phys.*, 10, 11707-  
813 11735.
- 814 Woodbury, P. B., et al. (2007), Carbon sequestration in the U.S. forest sector from 1990 to 2010,  
815 *Forest Ecology and Management*, 241, 14-27.
- 816 Yang, Z., et al. (2007), New constraints on Northern Hemisphere growing season net flux,  
817 *Geophysical Research Letters*, 34(12), L12807.
- 818 Zhang, Y.-C., et al. (2004), Calculation of radiative fluxes from the surface to top of atmosphere  
819 based on ISCCP and other global data sets: Refinements of the radiative transfer model and  
820 the input data, *Journal of Geophysical Research*, 109, D19105, doi:  
821 19110.11029/12003JD004457.
- 822 Zheng, D., et al. (2011), Carbon changes in conterminous US forests associated with growth and  
823 major disturbances: 1992-2001, *Environ. Res. Lett.*, 6, 014012.
- 824 Zhou, L., et al. (2003), Relation between interannual variations in satellite measures of northern  
825 forest greenness and climate between 1982 and 1999, *Journal of Geophysical Research-*  
826 *Atmospheres*, 108(D1).

827 Zhu, Z., and D. L. Evans (1994), US forest types and predicted percent forest cover from AVHRR  
828 data, *Photogrammetric Engineering and Remote Sensing*, 60, 525-531.

829  
830

831 **Figure Captions.**

832 **Figure 1.** Schematic diagram illustrating stock and flux (italicized) relationships between the  
833 forest sector and atmosphere. The entire forest sector net flux (sink) as defined by the stock-  
834 change approach is: Net Flux =  $\Delta C_{stocks} + \Delta C_{wood\ products}$ . Alternatively, using our model driven  
835 estimates of NEP it is: Net Flux =  $NEP - Wood\ Products\ Emissions - Fire$ .

836 **Figure 2.** Conterminous U.S. distribution of forest type groups shown with thick state boundaries  
837 that trace regions from the Resource Planning Act Assessment by the US Forest Service. Colors  
838 differentiate FIA forest type groups. The rectangles represent areas where gridded climate and  
839 phenology were used in the simulation of fluxes and stocks for each forest type within each  
840 rectangle.

841 **Figure 3.** Characteristic trajectories of aboveground live wood biomass regrowth and associated  
842 carbon sources / sinks (expressed as net ecosystem productivity, *NEP*) following a stand-  
843 replacing disturbance in high productivity Douglas-fir stands of the Pacific Northwest. Results  
844 are from the CASA model fit to regrow stocks consistent with 25 independent samples from the  
845 forest inventory data (red circles). Net releases in the year following disturbance are as low as -  
846  $3000\ g\ C\ m^{-2}\ a^{-1}$  (see Auxiliary Material 2, Figure A2.2) rising to above  $-500\ g\ C\ m^{-2}\ a^{-1}$  in the  
847 second year of regrowth.

848 **Figure 4a,b.** Map of average net ecosystem productivity (top, a) and uncertainty expressed as  
849 one standard deviation (bottom, b) ( $NEP$  in  $g\ C\ m^{-2}\ a^{-1}$ ) for forests of the conterminous US.

850

851

852 Table 1. Regional distribution of forest area, live biomass (Live  $B$ ), ratio of EPA [2008] to this  
 853 study's forest area ( $f_{\text{EPA08}}$  Area), ratio of EPA [2008] to this study's live biomass ( $f_{\text{EPA08}}$  Live  $B$ ),  
 854 net ecosystem productivity (NEP), fraction of forest that is less than 25 years old (<25y), less than  
 855 5 years old (<5y).

Region	Area [ $10^9 \text{ m}^2$ ]	Live $B$ [Tg C]	$f_{\text{EPA08}}$ Area [--]	$f_{\text{EPA08}}$ Live $B$ [-]	NEP [Tg C $\text{a}^{-1}$ ]	<25y [%]	<5y [%]
NE	339	3,253	1.11	1.01	$32 \pm 5.5$	10	2
NLS	212	1,236	0.99	1.11	$12 \pm 1.3$	16	3
SE	355	2,621	1.00	0.94	$30 \pm 3.5$	39	8
SC	384	3,220	1.27	1.00	$40 \pm 4.2$	37	8
RMN	192	1,189	0.98	1.10	$7 \pm 1.8$	21	5
RMS	493	1,815	0.81	0.97	$11 \pm 5.5$	1	0
PSW	127	1,522	1.06	0.95	$13 \pm 2.8$	11	2
PNW	202	2,162	1.05	1.13	$18 \pm 3.0$	19	4
Total/Mean	2,303	17,017	1.03	1.08	$164 \pm 27.7$	17	4

856  
 857

858 Table 2. Changes in carbon stocks [ $\text{Tg C a}^{-1}$ ] in the year 2005 reported in different studies.  
 859 Italicized values are inferred from mass balance.

	This Study	<i>EPA [2008]</i>
$\Delta\text{Total Soil C}$	3	9
$\Delta\text{Litter C}$	1	15
$\Delta\text{Coarse Woody Debris (CWD)}$	4	16
$\Delta\text{CWD Below}$	0	--
$\Delta\text{Live C}$	39	133
Total Stock Change	47	173
Removals <sup>g</sup>	117	<i>162</i>
Harvest <sup>a</sup>	<i>107</i>	132
Wildfire Emissions <sup>b</sup>	10	30
<i>NEP</i> <sup>c</sup>	164	335
Wood Products Emissions <sup>d</sup>	102	102
Wood Products Storage <sup>e</sup>	5	30
Forest Sector - Atmosphere Exchange <sup>f</sup>	52	203

<sup>a</sup>this study inferred as: Harvest = Removals – Wildfire Emissions;

<sup>b</sup>this study estimated wildfire emissions from the Global Fire Emissions Database v3 (GFED3) [van der Werf *et al.*, 2010];

<sup>c</sup>for the purposes of this table calculated as:

$NEP = \Delta\text{Total Soil C} + \Delta\text{Litter C} + \Delta\text{CWD} + \Delta\text{CWD Below} + \Delta\text{Live C} + \text{Removals}$ ; values differ from those in Table 1 due to differences in the method of aggregation and associated averaging of terms;

<sup>d</sup>this study adopted values reported in *EPA [2008]*;

<sup>e</sup>this study calculated as:

Wood Products Storage = Removals – Wood Products Emissions – Wildfire Emissions;

<sup>f</sup>this study calculated as:

Forest Sector-Atmosphere Exchange = *NEP* – Wood Products Emissions – Wildfire Emissions;

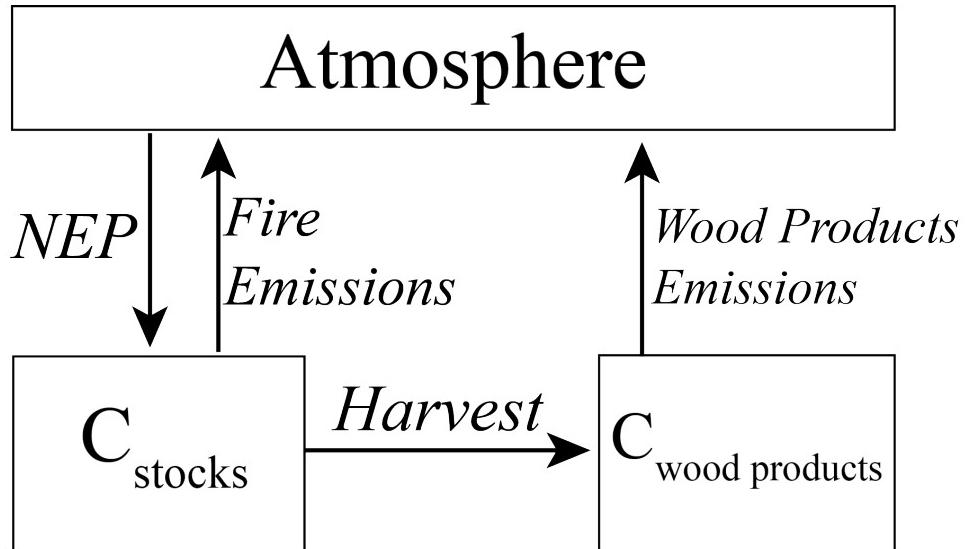
<sup>g</sup>for *EPA [2008]* calculated as: Removals = Wildfire Emissions + Harvest.

873 Table 3. Forest carbon *NEP* and stock change for the conterminous US [Tg C a<sup>-1</sup>] from this and a  
 874 sample of previously published estimates. Estimates are classified according to approach: age  
 875 structure–C accumulation (*A&A*), stock change ( $\Delta C_{stocks}=NBP$ ), or process model (*P*), where *P* is  
 876 a process model ensemble result that accounts for CO<sub>2</sub> and climate effects [Schimel *et al.*, 2000],  
 877 and Pacala *et al.* [2001] combines approaches for an overall estimate and range. Low and High  
 878 refers to 1 standard deviation about the mean estimate.

Source	Approach	Mean					
		<i>NEP</i>	Low	High	$\Delta C_{stocks}$	<i>Harvest</i>	<i>Fire</i>
<i>Schimel et al. [2000]<sup>a</sup></i>	<i>P</i>	80					
This Study <sup>b</sup>	<i>A&amp;A</i>	164	136	192	47	107	10
<i>Houghton et al. [1999]<sup>c</sup></i>	<i>A&amp;A</i>	182			10	92	80
<i>Turner [1995]<sup>e</sup></i>	<i>A&amp;A</i>	203			79	124	0
<i>Houghton [2003]<sup>d</sup></i>	<i>A&amp;A</i>	207			35	92	80
<i>Woodbury et al. [2007]<sup>f</sup></i>	$\Delta C_{stocks}$	270	256	293	108	132	30
<i>EPA [2008]<sup>g</sup></i>	$\Delta C_{stocks}$	335			173	132	30
<i>Birdsey &amp; Heath [1995]<sup>h</sup></i>	$\Delta C_{stocks}$	368			211	127	30
<i>Hurtt et al. [2002]<sup>i</sup></i>	<i>A&amp;A</i>	372	282	442	230	92	50
<i>Pacala et al. [2001]<sup>j</sup></i>	synthesis	392	312	472	220	92	80
<i>King et al. [2007]<sup>k</sup></i>	$\Delta C_{stocks}$	411	383	439	236	145	30

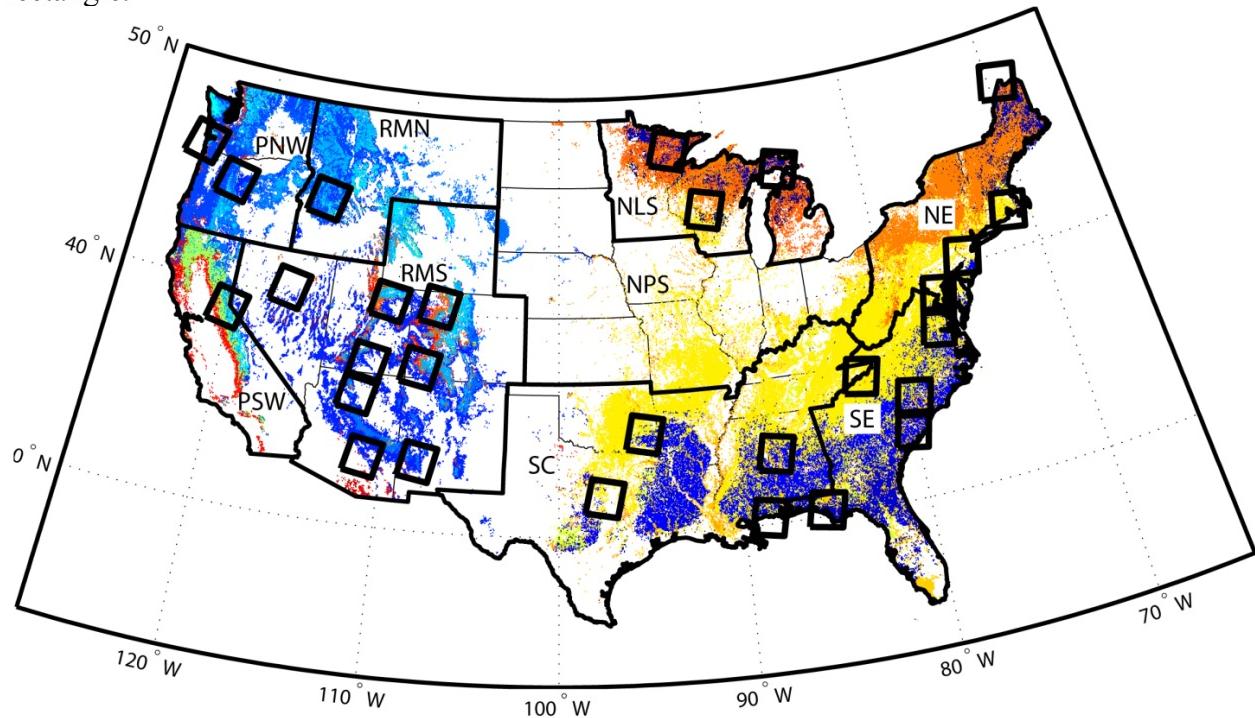
- 879 a) for 1980-1993
- 880 b) for 2005, C stock change = *NEP* – *Harvest* – *Fire*, see Table 2, our total removals are 117
- 881 Tg C a<sup>-1</sup> that includes fire and harvest, assume fire at 10 Tg C a<sup>-1</sup> (see GFED3 of van der
- 882 Werf *et al.*[2010] and Zheng *et al.* [2011])
- 883 c) 1980s
- 884 d) for 1990's, harvest and fire from Houghton *et al.* [1999]
- 885 e) for ~1990,
- 886 f) for 2005
- 887 g) for 2005
- 888 h) for 1992
- 889 i) for 1980s
- 890 j) for 1980's
- 891 k) for 1980s

893      Figure 1. Schematic diagram illustrating stock and flux (italicized) relationships between the  
894      forest sector and atmosphere. The entire forest sector net flux (sink) as defined by the stock-  
895      change approach is: Net Flux =  $\Delta C_{stocks} + \Delta C_{wood\ products}$ . Alternatively, using our model driven  
896      estimates of NEP it is: Net Flux =  $NEP - Wood\ Products\ Emissions - Fire\ Emissions$ .  
897



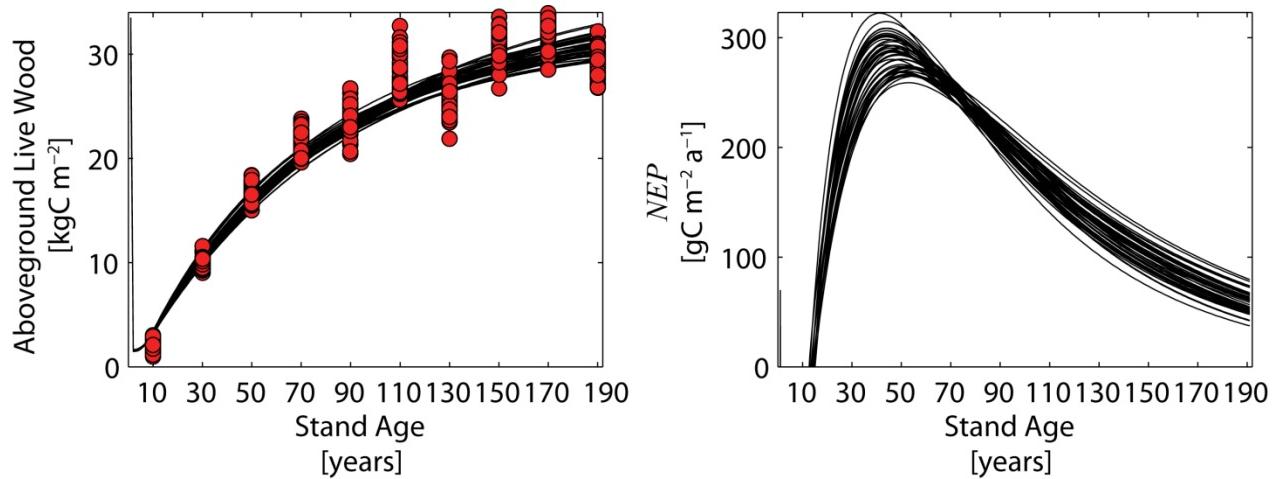
898  
899  
900

901      Figure 2. Conterminous U.S. Forest Type Groups shown with thick state boundaries that trace  
902      regions from the Resource Planning Act Assessment by the US Forest Service. Colors  
903      differentiate FIA forest type groups. The rectangles represent areas where gridded climate and  
904      phenology were used in the simulation of fluxes and stocks for each forest type within each  
905      rectangle.



906  
907  
908  
909  
910  
911

912     Figure 3. Characteristic trajectories of aboveground live wood biomass regrowth and associated  
913     carbon sources / sinks (expressed as net ecosystem productivity,  $NEP$ ) following a stand-  
914     replacing disturbance in high productivity Douglas-fir stands of the Pacific Northwest. Results  
915     are from the CASA model fit to regrow stocks consistent with 25 independent samples from the  
916     forest inventory data (red circles). Net releases in the year following disturbance are as low as -  
917      $3000 \text{ g C m}^{-2} \text{ a}^{-1}$  (see Auxiliary Material 2, Figure A2.2) rising to above  $-500 \text{ g C m}^{-2} \text{ a}^{-1}$  in the  
918     second year of regrowth.  
919



920

Figure 4a,b. Map of average net ecosystem productivity (top, a) and uncertainty expressed as one standard deviation (bottom, b) ( $NEP$  in  $\text{g C m}^{-2} \text{ a}^{-1}$ ) for forests of the conterminous US.

